

Petrology, mineralogy and geochemistry of olivine-phyric shergottites.

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Olivine-phyric shergottites are characterized by olivine megacrysts and a fine-grained matrix, recording a sequence of igneous processes, including the evolution of Martian mantle-derived magmas and crystallization within magma chambers [1]. In this study, we conduct petrological, mineralogical, and geochemical analyses of an olivine-phyric shergottite, Amgala 001, to understand its crystallization sequence in the Martian interior and to discuss the magmatic activity. For comparison, other olivine-phyric shergottites, Northwest Africa (NWA) 1195, NWA 2626, and Tissint are also examined. Polished mounts of Amgala 001 and three other samples were briefly observed using a scanning electron microscopy with energy dispersive X-ray spectroscopy (SEM-EDS; JEOL JSM-6390A) at Hiroshima University (HU). Detailed mineralogy were observed using a field emission SEM-EDS combined with a focused ion beam (FIB) system (Thermo Fisher Scientific Helios G4 UC) and a transmission electron microscopy (TEM; JEOL JEM-F200) at the Natural Science Center for Basic Research and Development (N-BARD), HU. Major element compositions of individual minerals were analyzed using an electron probe microanalyzer (EPMA; JEOL JXA-iSP 100) at N-BARD, HU. Rare earth element (REE) abundances of whole-rock sample of Amgala 001 were determined by an inductively coupled plasma mass spectrometry (ICP-MS) at the Kochi Institute for Core Sample Research, Japan Agency for Marine-Earth Science and Technology (JAMSTEC). Major element compositions of the same whole rock sample were determined using an inductively coupled plasma optical emission spectroscopy (ICP-OES) at HU.

The whole-rock analysis of Amgala 001 shows a flat REE pattern with the La/Yb ratio of 1.09, indicating the meteorite is classified as enriched shergottite. The SEM-EDS and EPMA analyses reveal that most olivine crystals in Amgala 001 exhibit a well-developed positive zoning, with Mg-rich cores (average Mg# = 71) and Fe-rich rims (average Mg# = 51). The Mg# value of coexisting melt is calculated as 56, using an Fe^{2+} -Mg partition coefficient between olivine and the Martian magma [2]. This value is much lower than the analyzed whole-rock Mg content (Mg# = 56.4), indicating that the bulk compositions of Amgala 001 reflect the olivine accumulation.

Elemental mapping of an olivine grain in Amgala 001 reveals oscillatory phosphorus (P) band zoning toward the Fe-rich rim. The submicron-scale analyses using scanning transmission electron microscopy (STEM) and EDS demonstrate that the P-band zoning is not associated with mineral inclusions, but instead reflects chemical heterogeneities within olivine crystals. This P-band zoning is interpreted to record rapid crystal growth of olivine [3, 4].

Crystallization pressure of the pyroxene can be estimated based on their Ti and Al contents and an experimentally derived model by [5] as 0.93 GPa, corresponding to the lower crust of Mars. At this depth, cores of olivine megacrysts, pyroxene, and Fe-Ti-Al oxides crystallized in a magma chamber.

Subsequently, plagioclase crystallized within the magma chamber or during the magma ascent. The occurrence of P-band zoning at the center of the olivine grain indicates that the glomerocrystic olivine accumulated within a second-stage magma chamber under shallower and lower-temperature conditions. These crystals were finally entrained during magma ascent and erupted onto the Martian surface.

References:

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- [2] Filiberto and Dasgupta (2011), *Earth Planet. Sci. Lett.*, 304, 527–537.
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[5] Filiberto et al. (2010), *Meteorit. Planet. Sci.*, 45, 1258–1270.

Keywords: Martian meteorites, olivine-phyric shergottites, Magmatic processes, Phosphorus zoning, Amgala 001

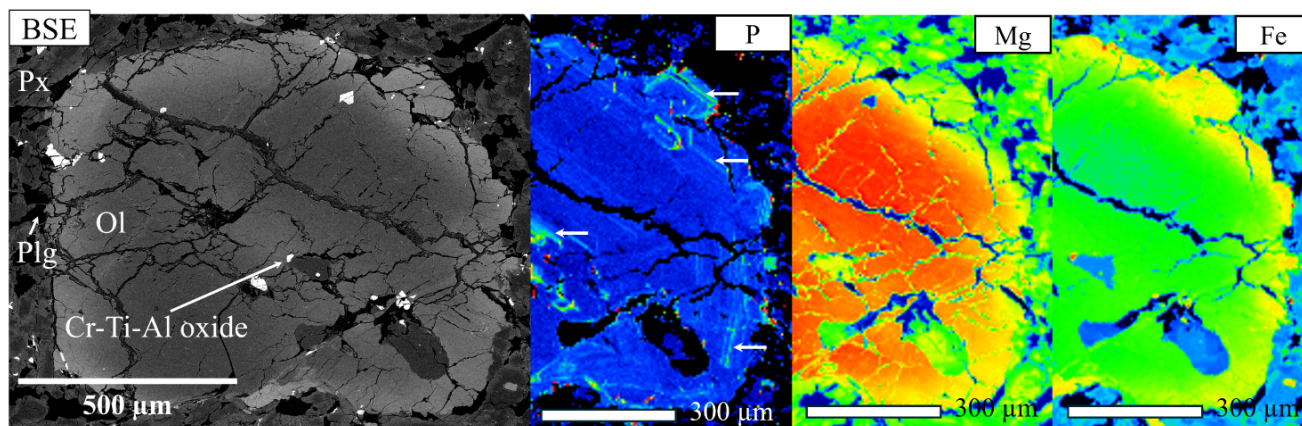


Figure 1. Large-sized olivine in Amgala 001: BSE image and K α X-ray elemental heat maps. P K α map after masking pixels with high Ca K α intensities.